



ERDC
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Long-term Morphologic Modeling at Coastal Inlets

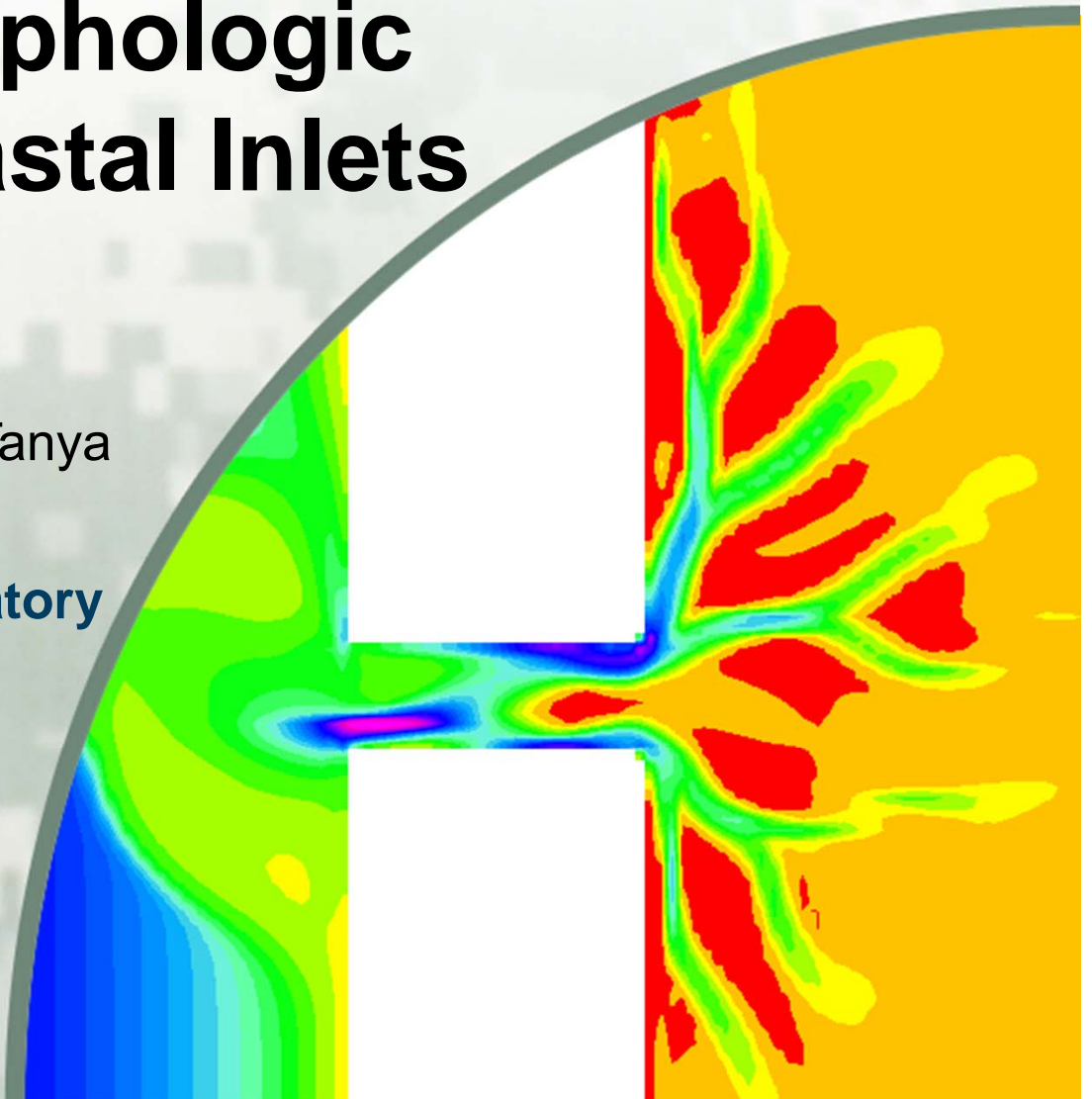
Alex Sánchez

Richard Styles, Mitchell Brown, Tanya
Beck, and Honghai Li

Coastal and Hydraulics Laboratory
US Army Corps of Engineers



**US Army Corps
of Engineers®**

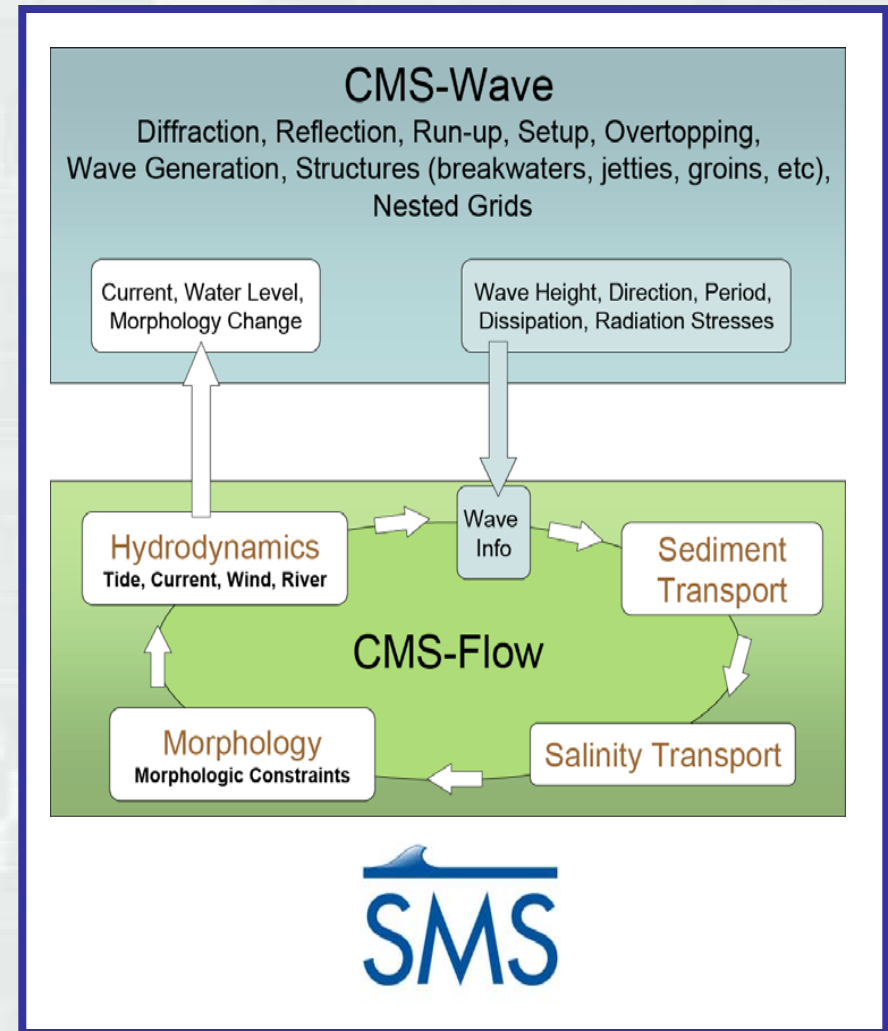


Introduction

- Motivation:
 - ▶ Prediction of morphodynamic processes at coastal inlets is challenging but crucial for coastal sediment management, navigation, channel maintenance, and breach erosion protection
- Issue:
 - ▶ Difficult to conduct meaningful long-term validation of morphodynamic models using real data
- Approach:
 - ▶ Simulate idealized inlets representing 9 US inlets and compare inlet evolution, characteristics, and features with the actual inlets empirical formulas (soft validation)

Introduction: Coastal Modeling System

- Hydrodynamics:
 - ▶ 2DH shallow-water equations
 - ▶ Fully implicit, finite-volume method
 - ▶ Non-uniform or Telescoping Cartesian grids
- Sediment Transport
 - ▶ Inline
 - ▶ Total-load non-equilibrium sediment transport
 - ▶ Erosion/deposition calculated using an adaptation approach
 - ▶ Several options for transport capacity formula
- Waves
 - ▶ Spectral wave-action balance equation
 - ▶ Implicit finite-difference method



Empirical Relations

■ Cross-sectional area

- ▶ O'Brien (1931, 1969), Kraus (1998), Jarrett (1976), van der Kreeke (1992), Powell et al. (2006), etc.

$$A = CP^n$$

$A \rightarrow$ Cross-sectional area [m^2]

■ Ebb tidal shoal volume

- ▶ Walton and Adams (1976)

$$V_{ebb} = aP^b$$

$P \rightarrow$ Tidal prism [m^3]

$C \rightarrow 8.83 \times 10^{-6} - 1.88 \times 10^{-3} [m^{-1}]$

$n \rightarrow 0.81 - 1.10 [-]$

- ▶ Hicks and Hume (1996)

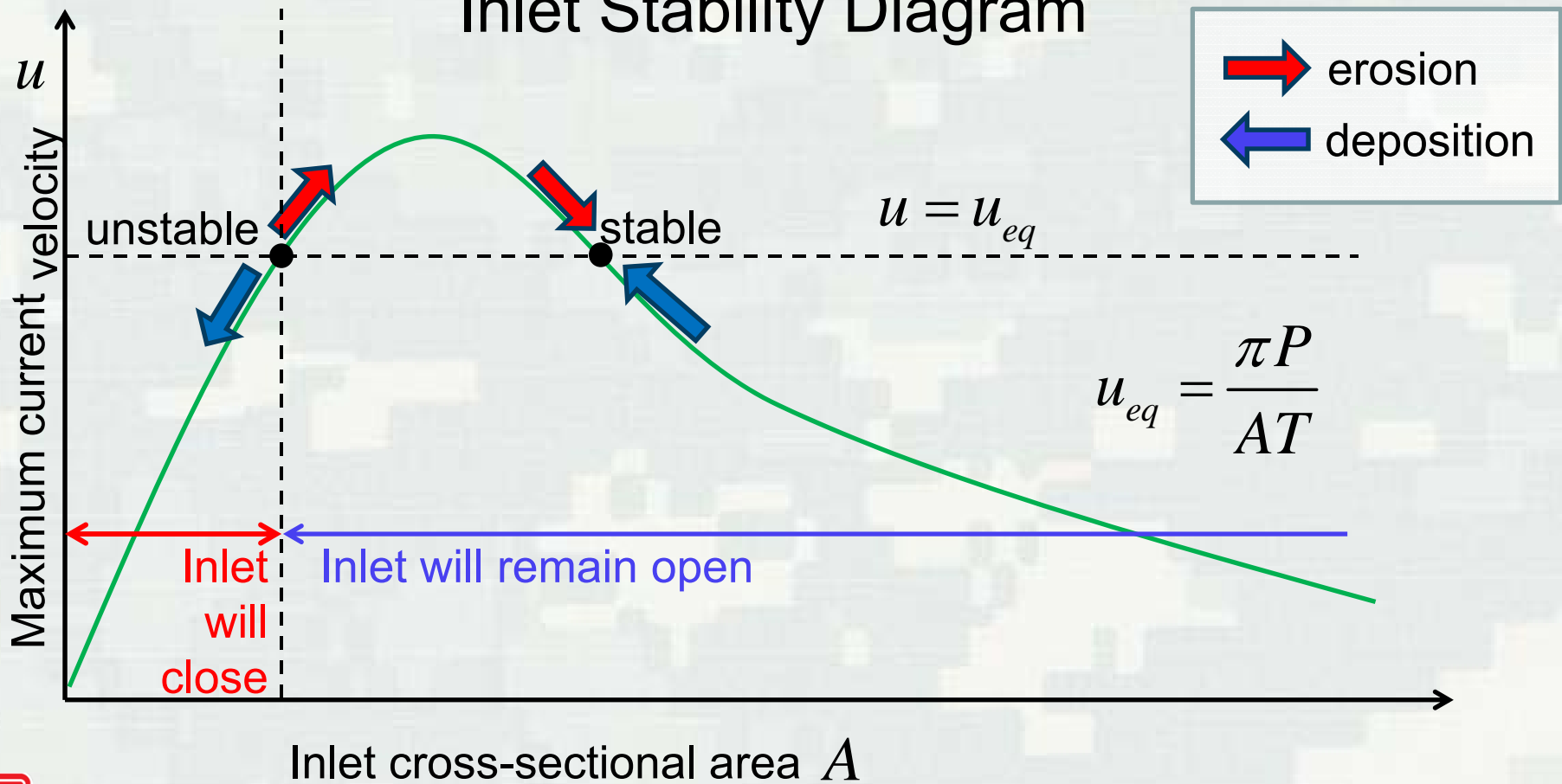
$$V_{ebb} = 1.37 \times 10^{-3} P^{1.32} (\sin \theta)^{1.33}$$

$a \rightarrow 5.3 \times 10^{-3} - 8.4 \times 10^{-3}$

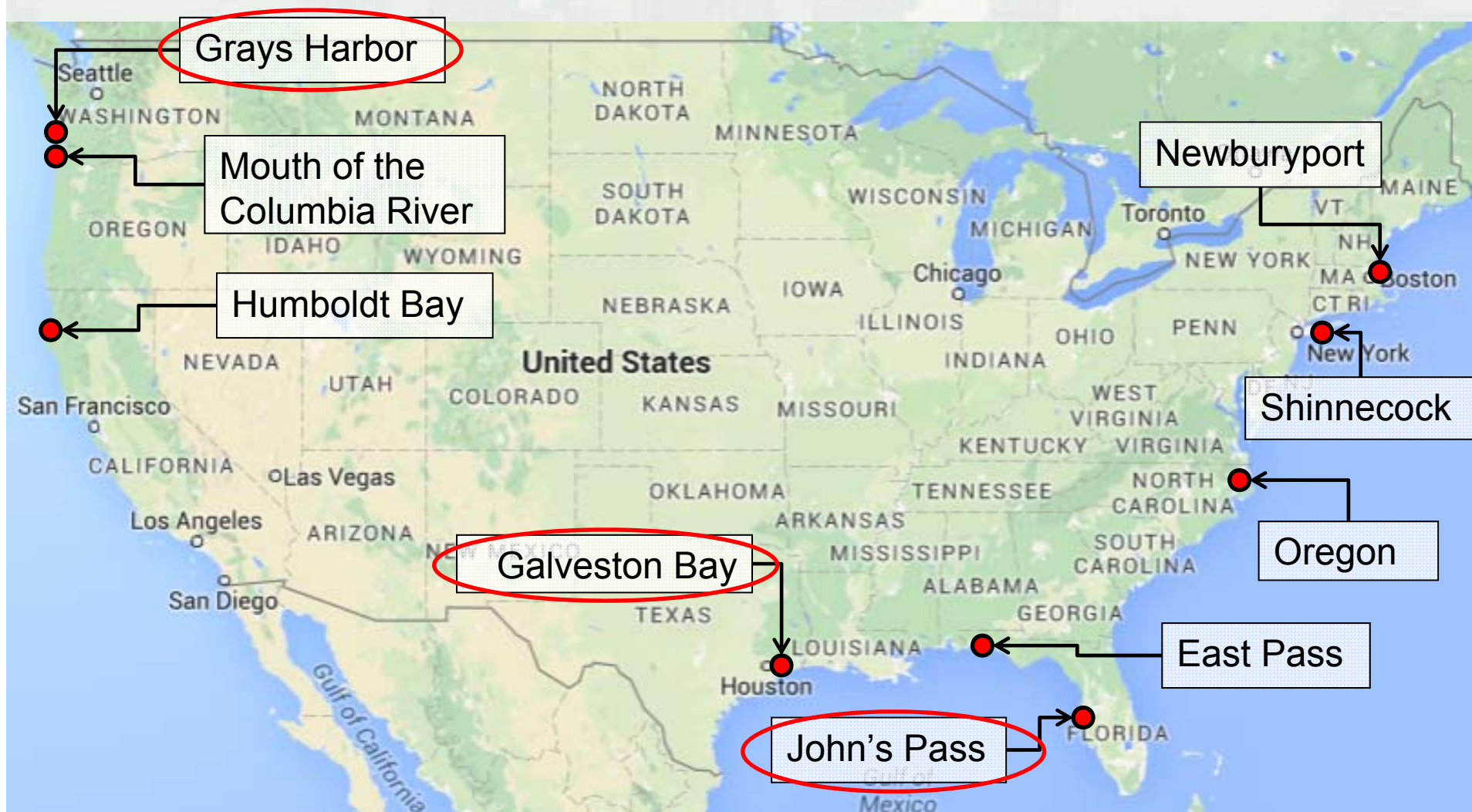
$b \rightarrow 1.23$

Inlet Stability Analysis

Escoffier's (1940) Inlet Stability Diagram

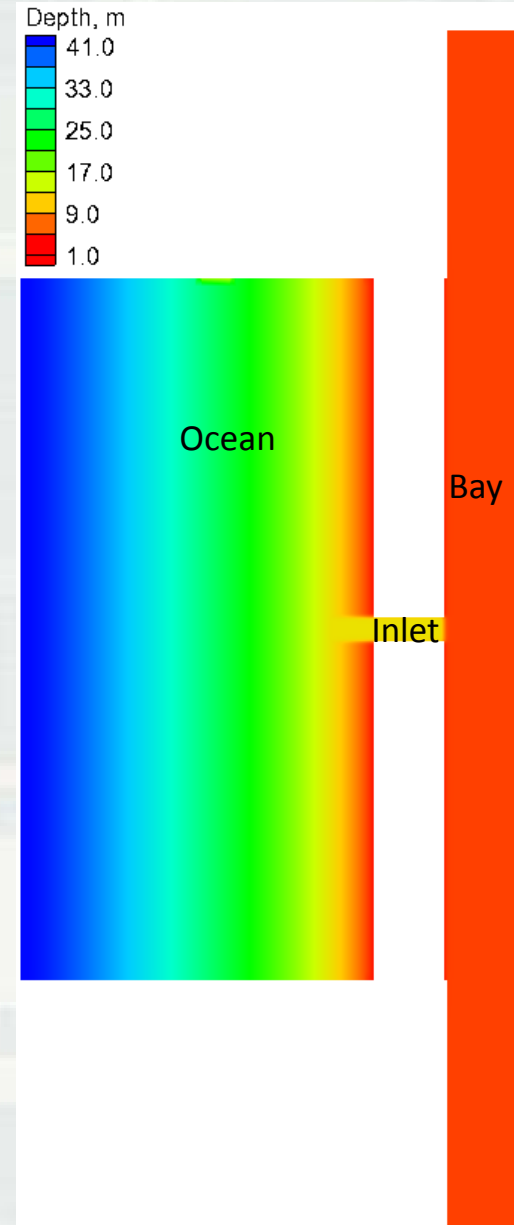


Methods: Base Inlets



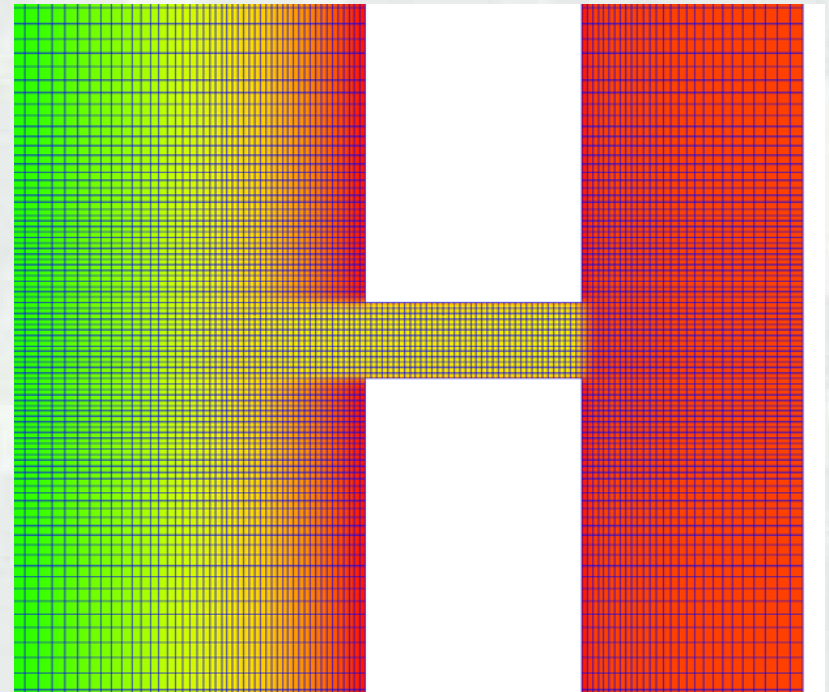
Methods: Idealized Inlets

- Initial Morphology
 - ▶ Equilibrium offshore profile based on measured bathymetry or median grain size
 - ▶ Flat rectangular bay with dimensions based on actual inlet. Bay width and length adjusted to match actual bay area
 - ▶ Flat rectangular inlet with width and area matching actual inlet
- Water levels
 - ▶ Tidal constituents
- Waves
 - ▶ Representative year based on mean sediment transport rate estimated from the CERC formula and nearby buoy data



Methods: Model Setup

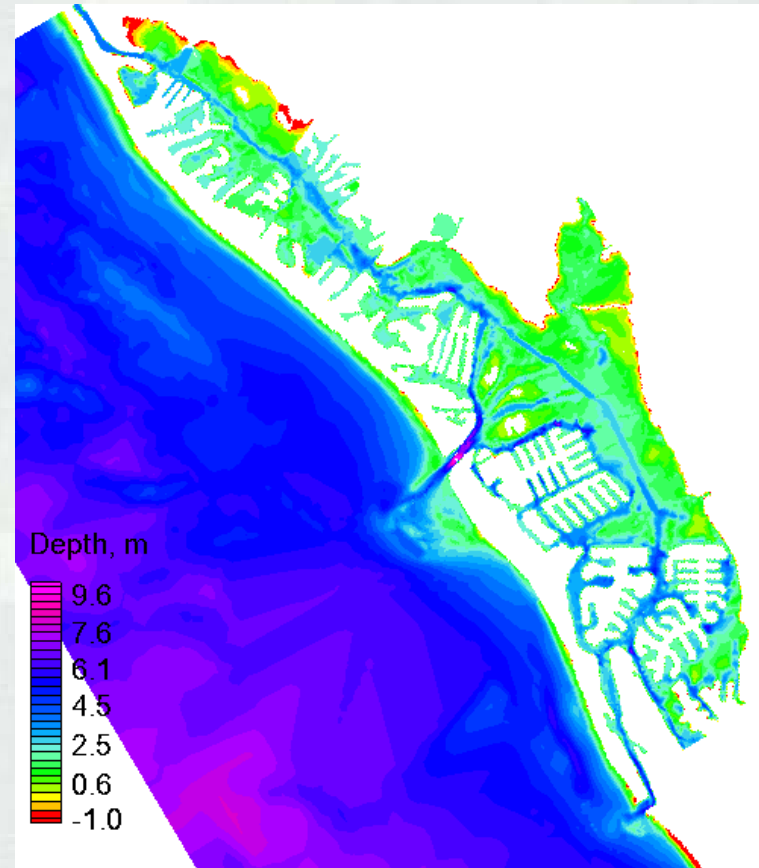
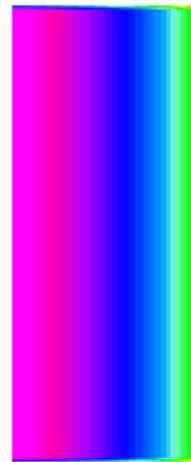
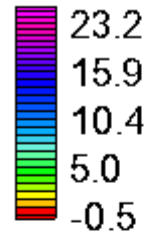
- Flow
 - ▶ Manning's $n = 0.025 \text{ s/m}^{1/3}$
 - ▶ Coriolis
- Sediment transport
 - ▶ Single representative grain size
 - ▶ Morphologic acceleration factor = 10
- Time stepping
 - ▶ Flow and sediment: 15 min
 - Second-order scheme
 - ▶ Waves: 1 hr
- Grids
 - ▶ Same for flow, sediment, and waves
 - ▶ Resolution
 - At least 10 cells across inlet



John's Pass, FL

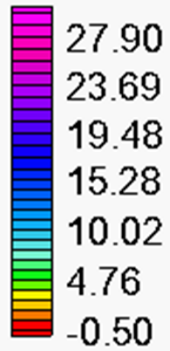
- Waves
 - ▶ $H_{mo} = 0.73$ m
 - ▶ $T_p = 4$ s
- Tidal range
 - ▶ 0.43 m
- Bay Dimensions
 - ▶ Area = $4.5e7$ m²
 - ▶ Length = 27 km
 - ▶ Width = 19 km
- Inlet
 - ▶ Area = 845 m²
 - ▶ Width = 300 m

Depth, m



Johns Pass, FL

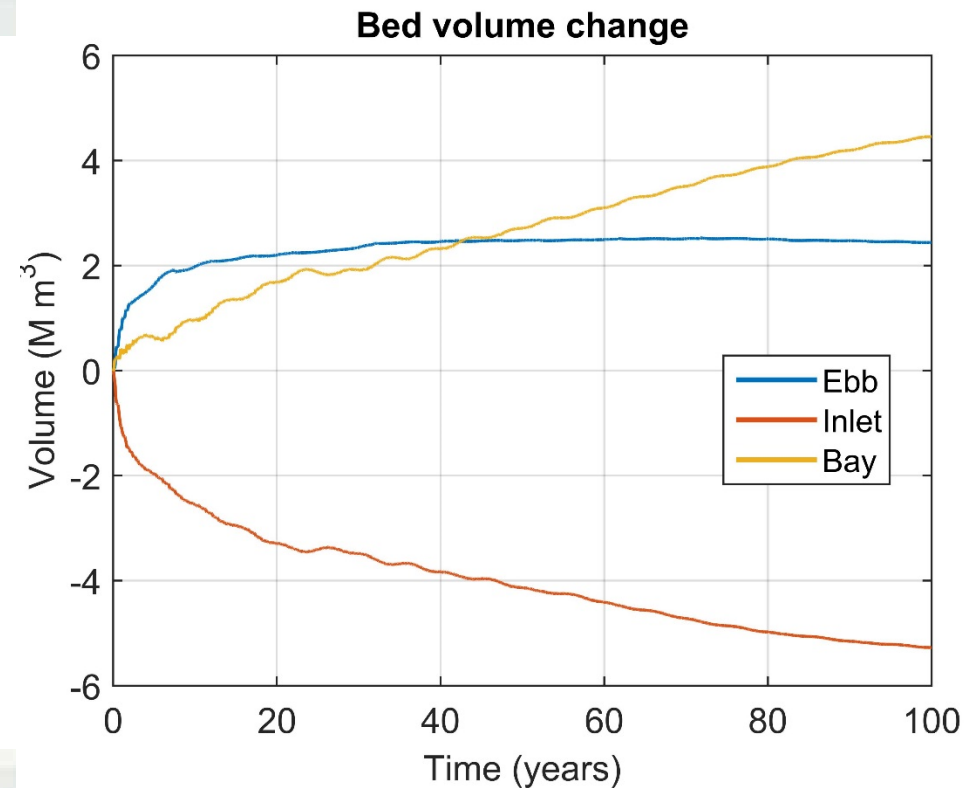
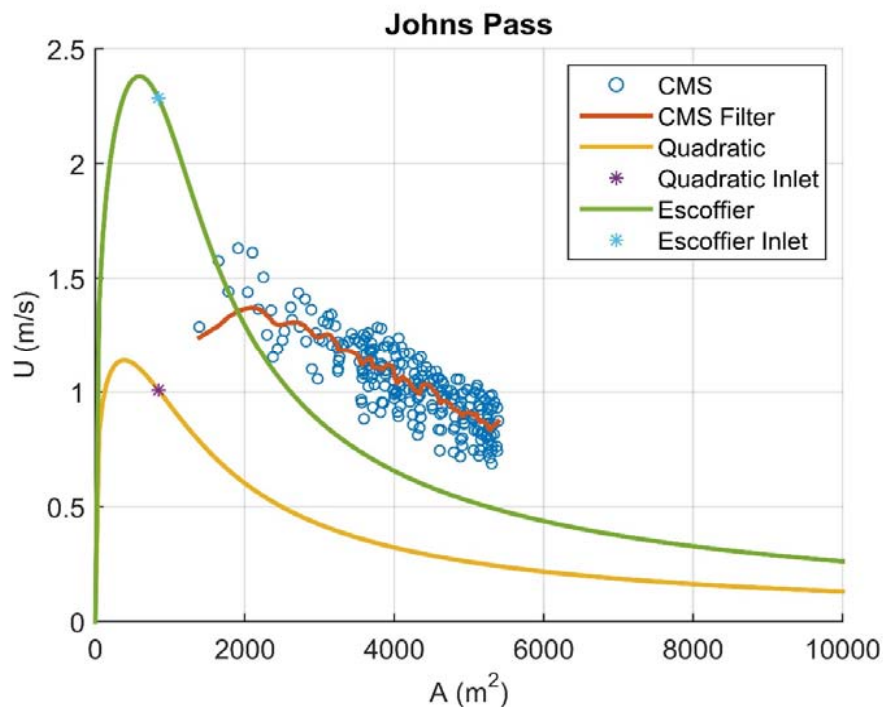
Depth, m



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Results: Johns Pass, FL

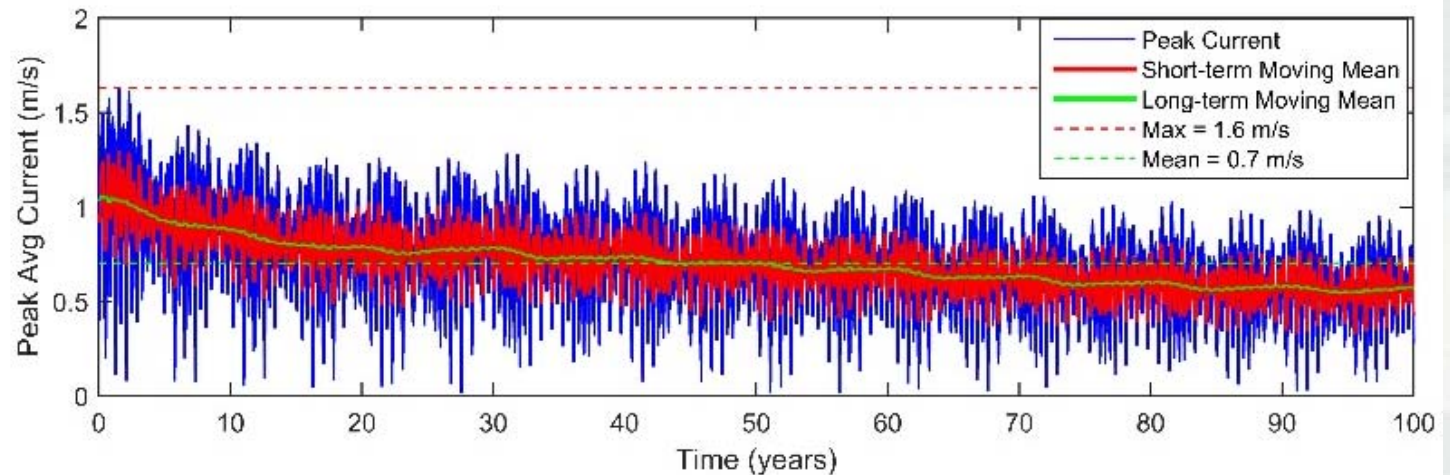
Flood dominant



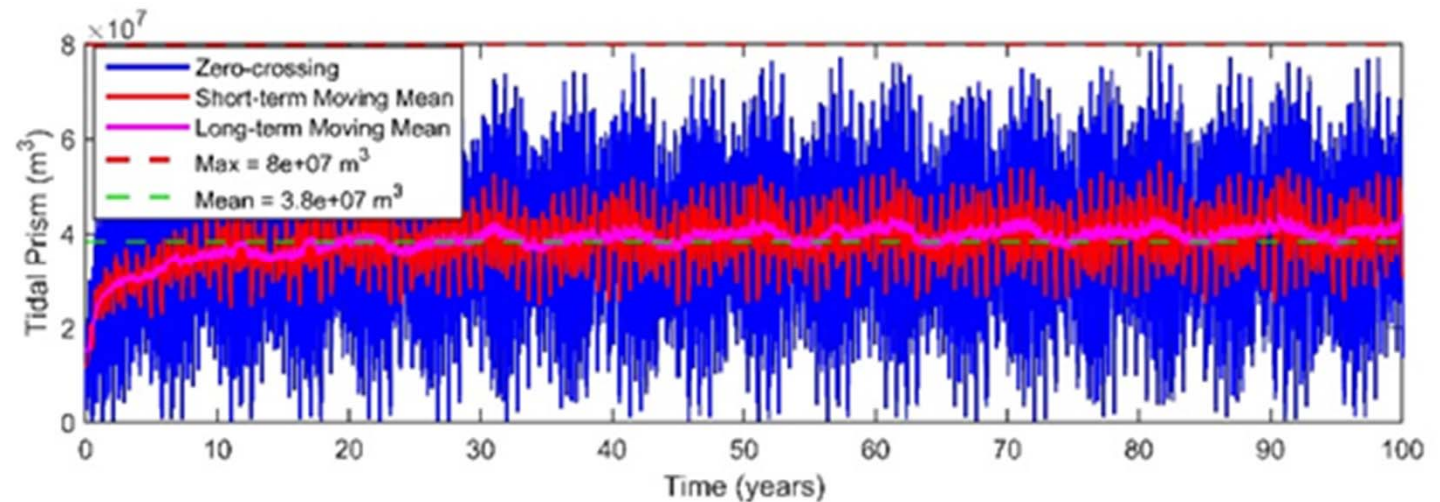
- Actual ebb shoal volume
 - 2.1 to 2.3 M m³

Results: Johns Pass, FL

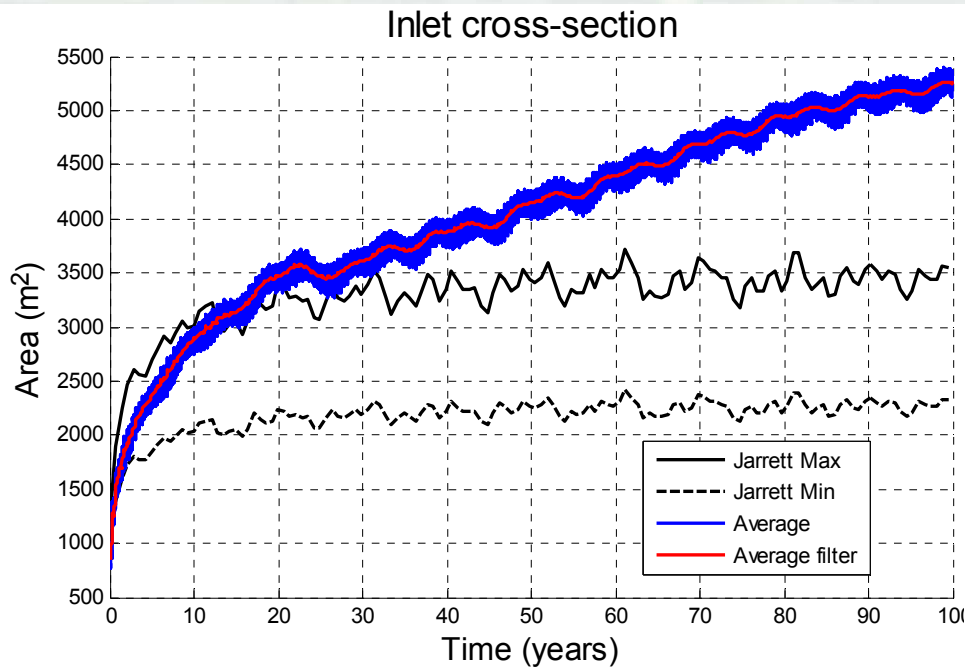
Actual peak
current
velocity
~1.2 m/s



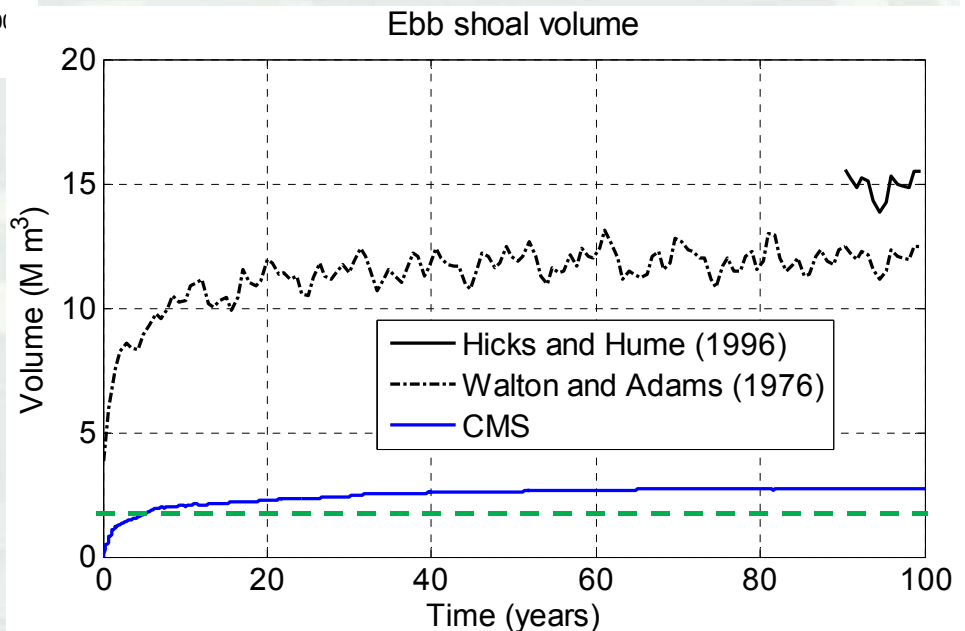
Tidal Prism:
 $2.1 \times 10^7 \text{ m}^3$



Results: Johns Pass, FL



- Inlet does not reach equilibrium
- Ebb shoal does reach equilibrium but is underestimated



Results: Grays Harbor

Waves

$$H_{mo} = 2 \text{ m}$$

$$T_p = 10 \text{ s}$$

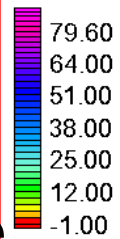
Tidal range

2.25 m

Inlet

$$A_c = 31200 \text{ m}^2$$

Depth, m



Initial
bathymetry

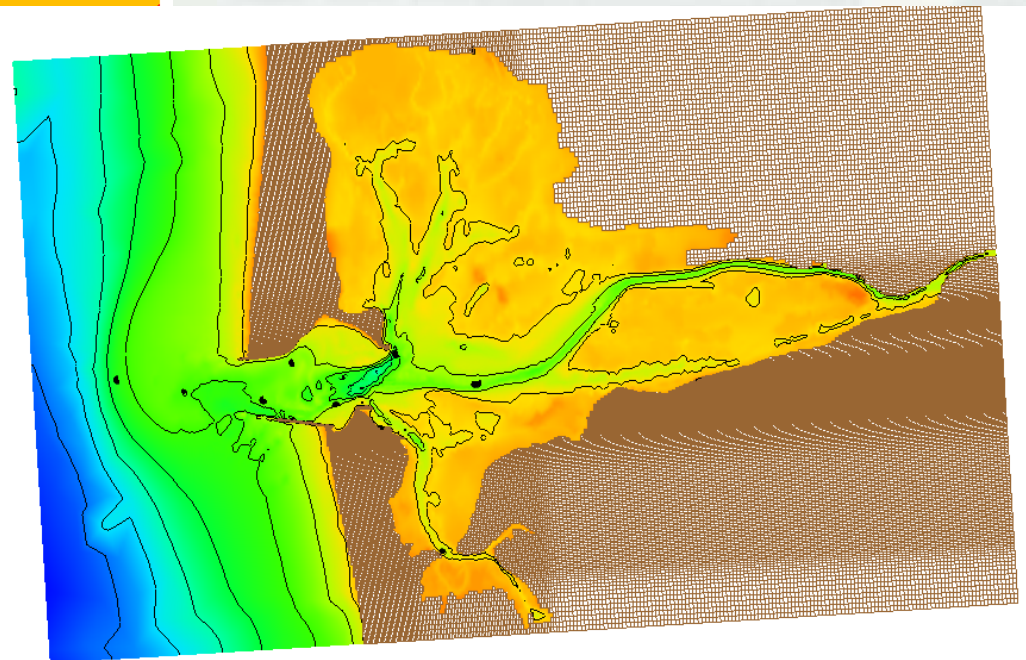
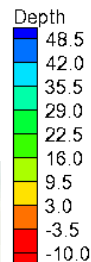
Bay

$$A_b = 513 \text{ M m}^2$$

$$W = 19 \text{ km}$$

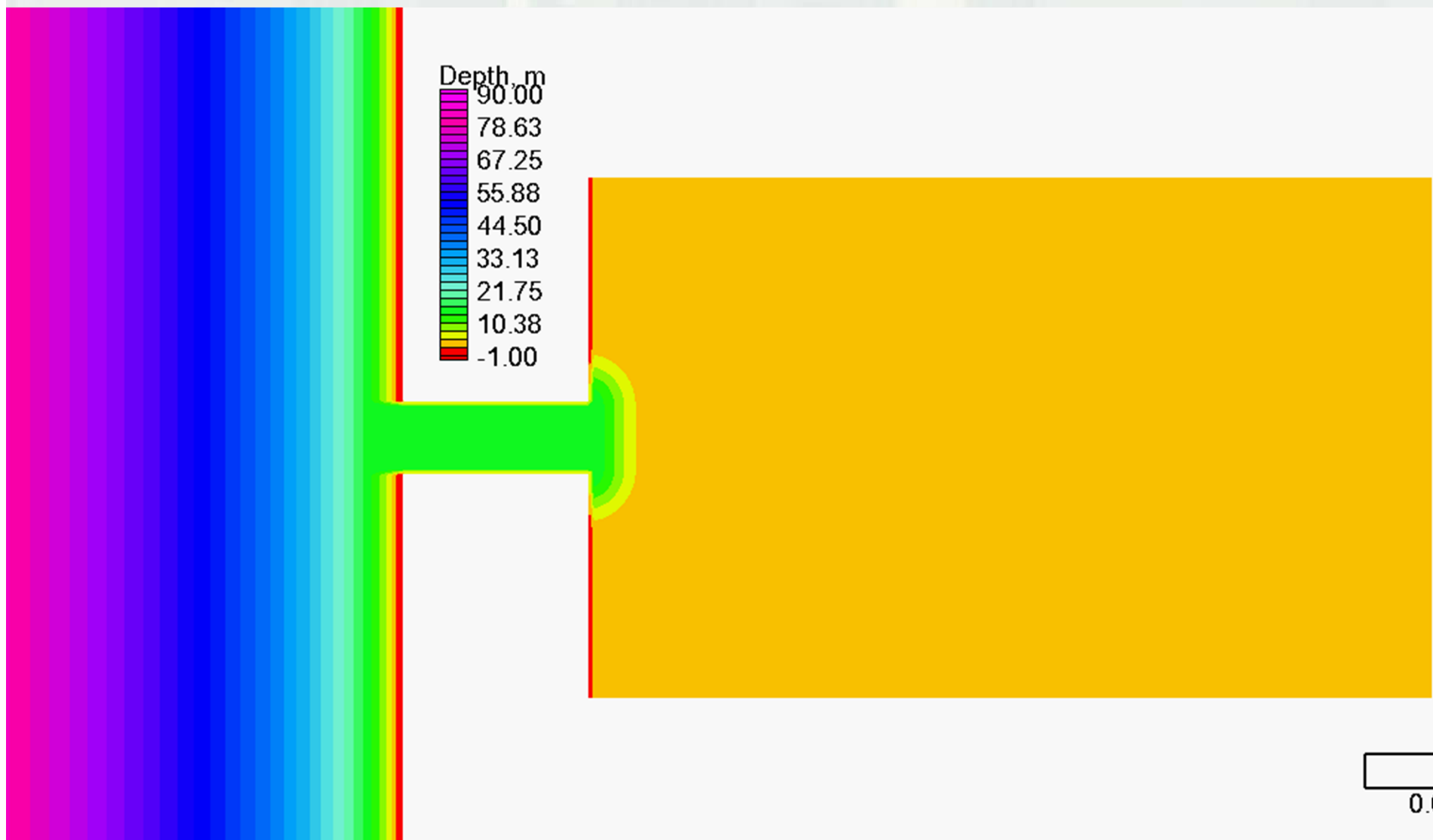
$$L = 27 \text{ km}$$

Actual bathymetry

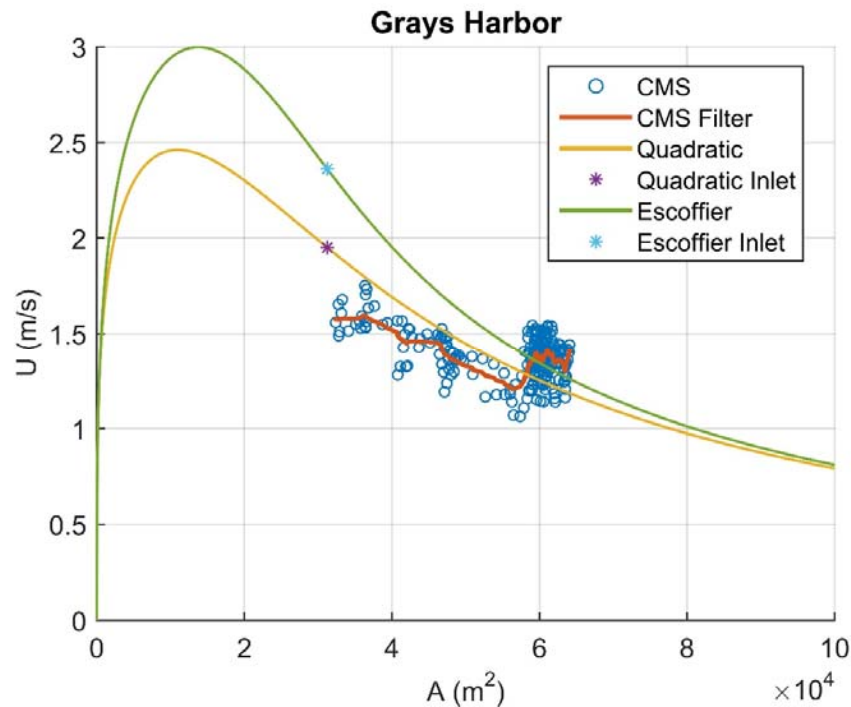


BUILDING STRONG®

Results: Grays Harbor, WA

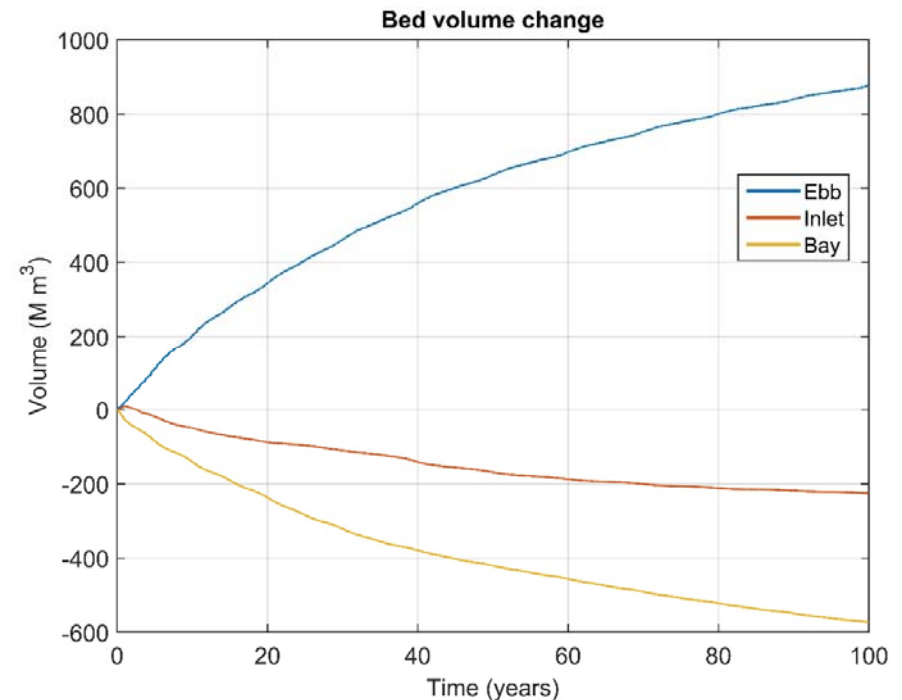


Grays Harbor, WA

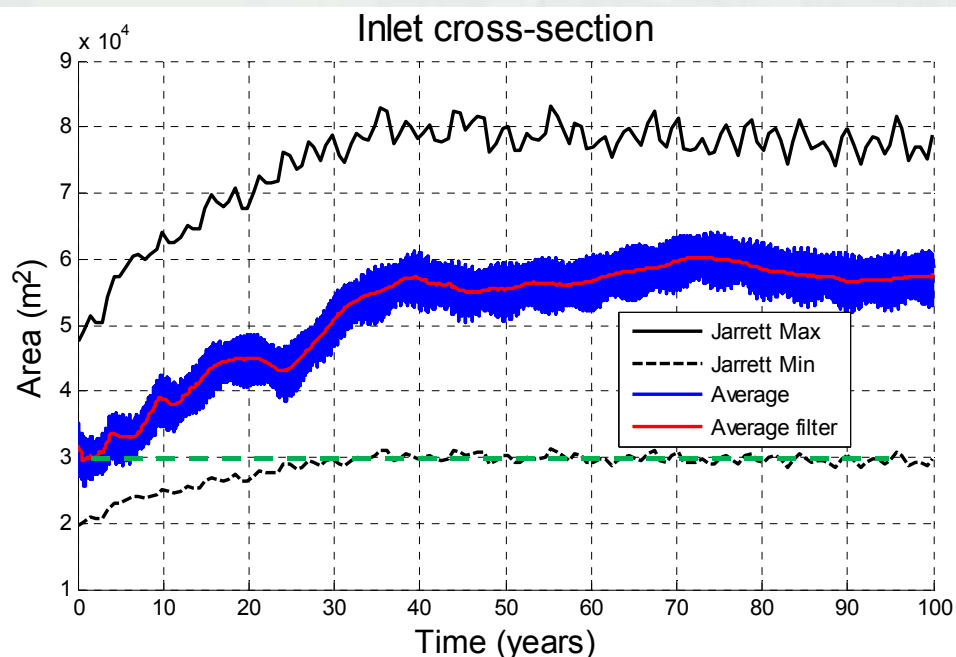


Equilibrium cross-sectional area of idealized inlet larger than initial condition

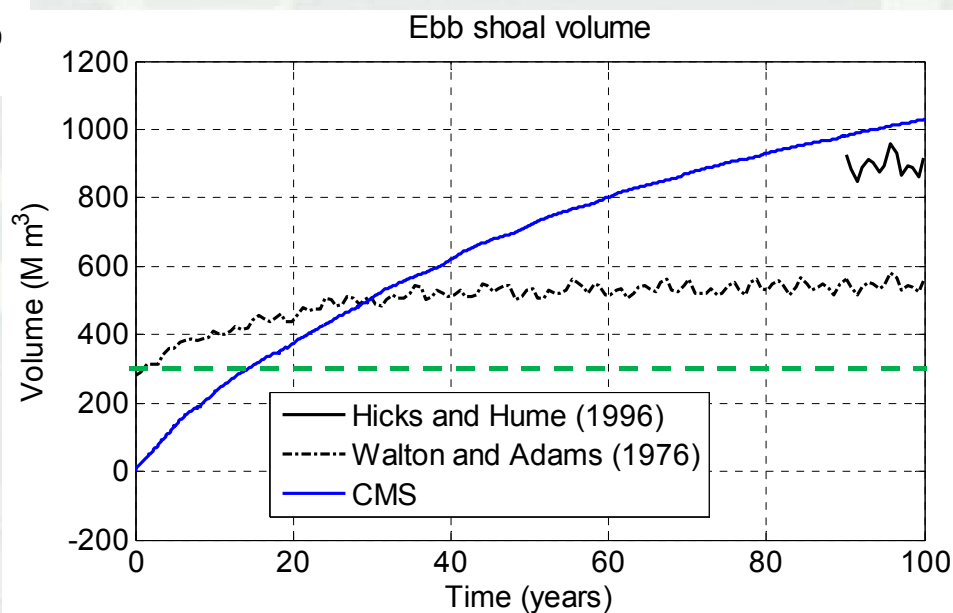
Inlet still evolving after 100 years



Results: Grays Harbor, WA



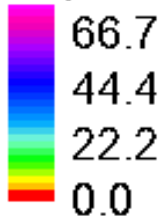
- Actual ebb shoal volume
► 240 to 250 M m^3



Galveston, TX

Initial bathymetry

Depth, m



Waves

$$H_{mo} = 1.2 \text{ m}$$

$$T_p = 5 \text{ s}$$

Tidal range
0.43 m

Inlet

$$A_c = 16800 \text{ m}^2$$

$$W = 3 \text{ km}$$

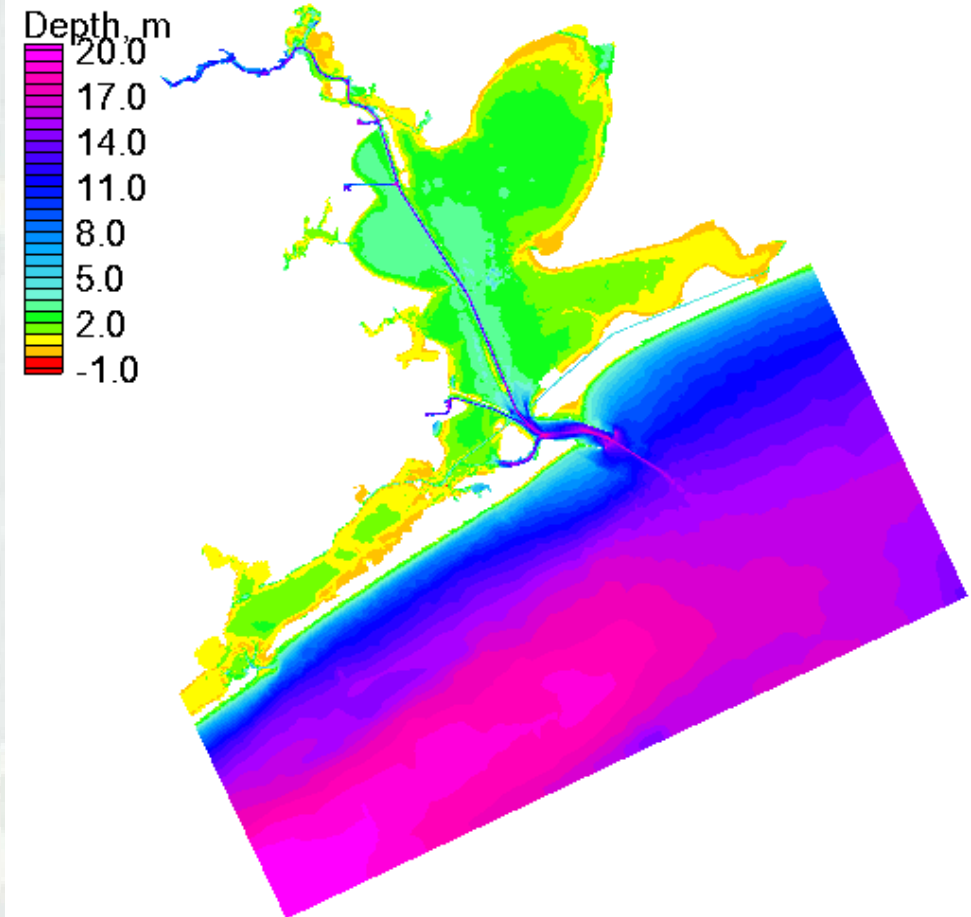
$$L = 7.5 \text{ km}$$

Bay

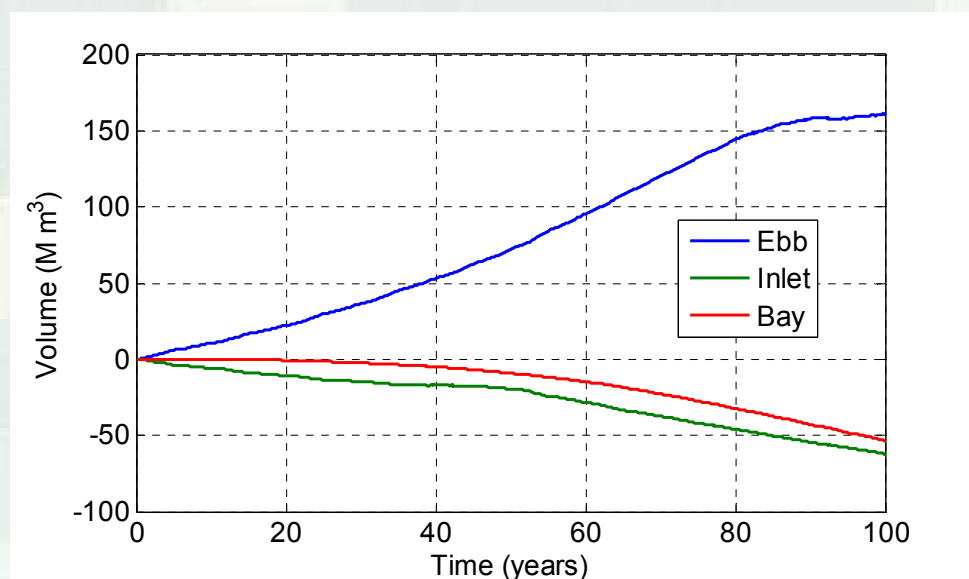
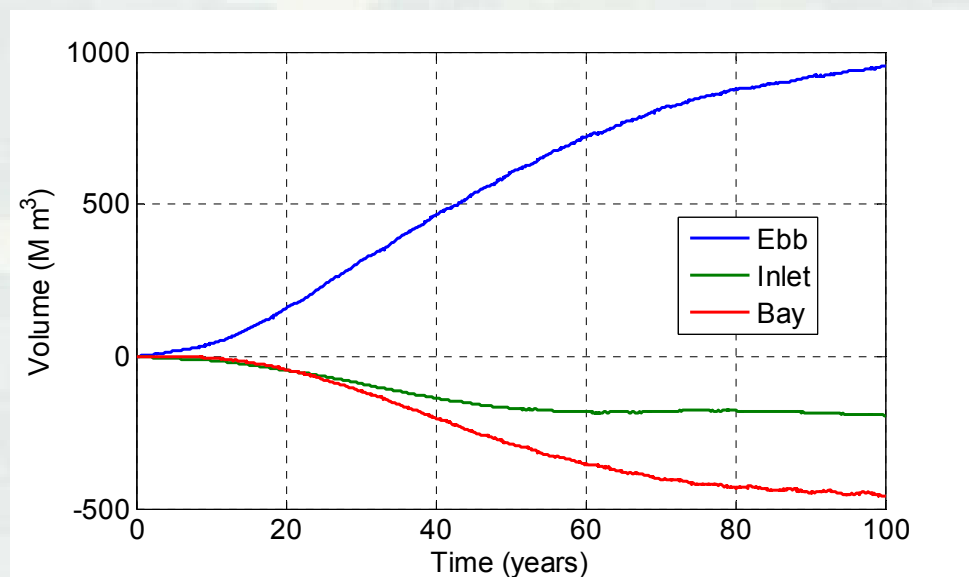
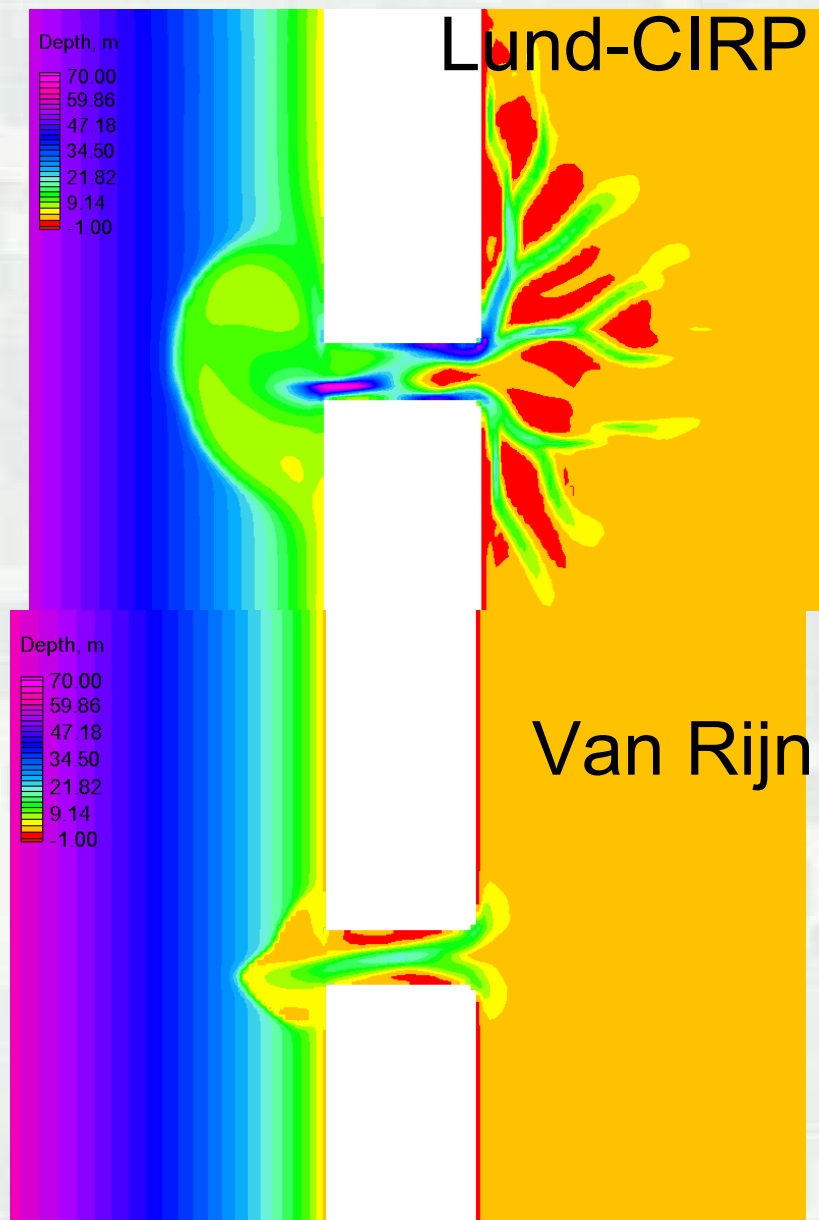
$$A_b = 1600 \text{ M m}^2$$

$$W = 50 \text{ km}$$

$$L = 32 \text{ km}$$

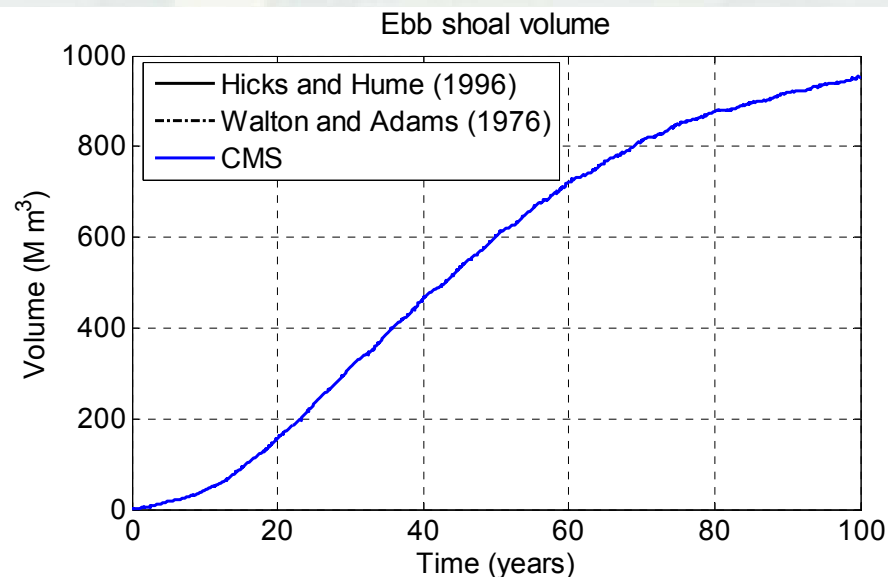
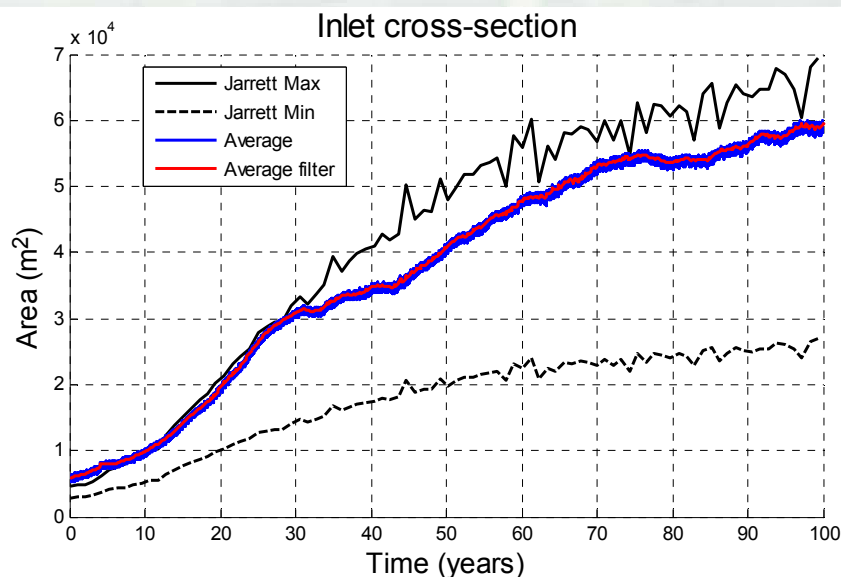


Results: Galveston, TX

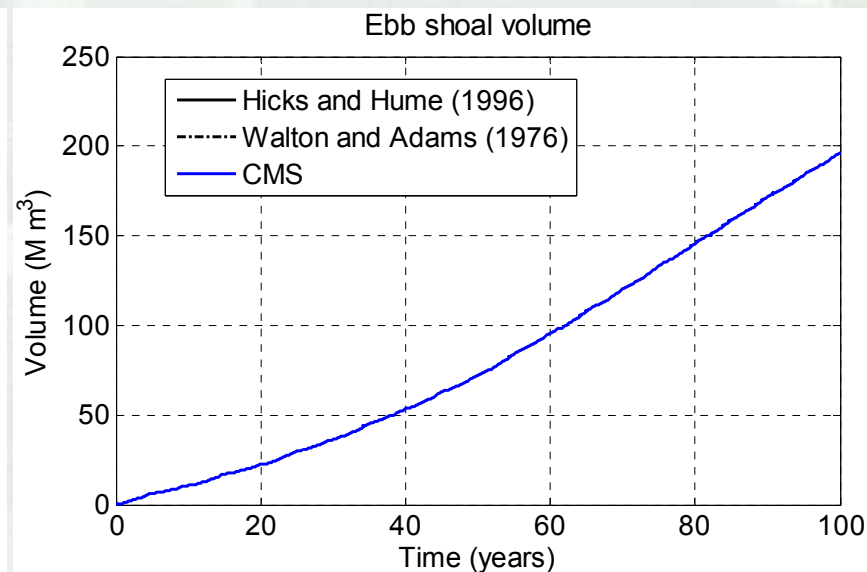
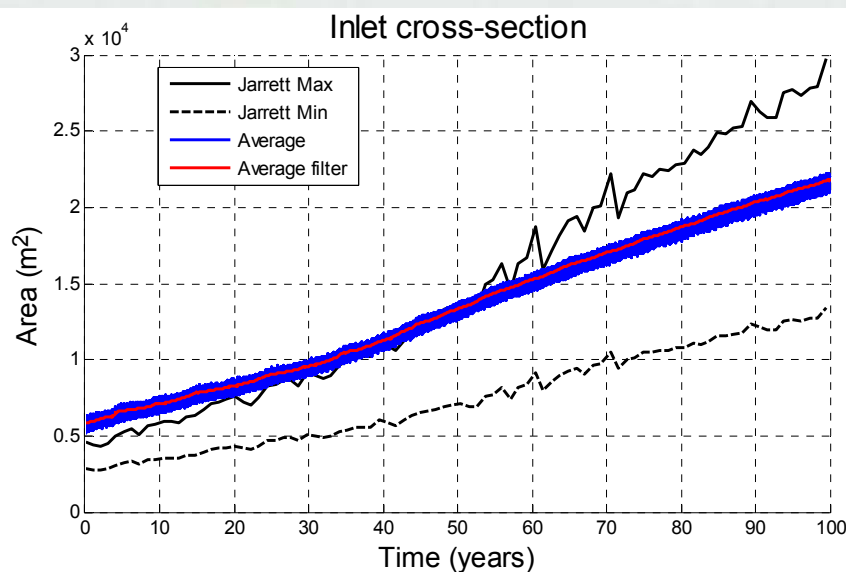


Results: Galveston, TX

Lund-CIRP



Van Rijn



Discussion and Conclusions

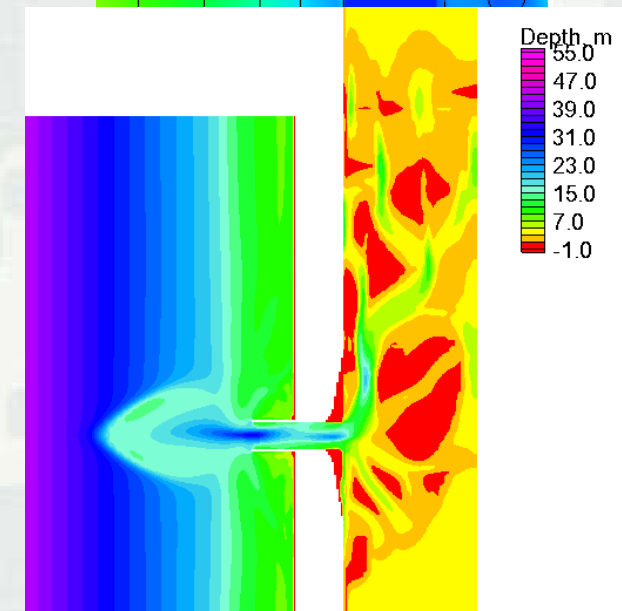
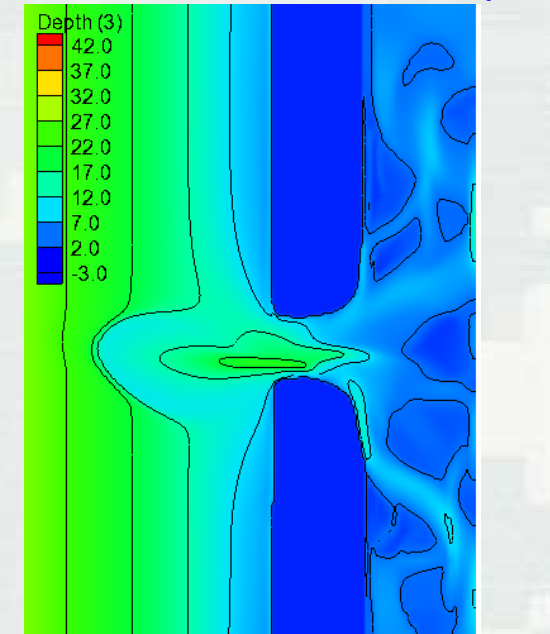
- Rate of bed change within the first 10-20 years is rapid and then slows
- None of the simulated inlets reached a full dynamic equilibrium after 100 years suggesting that either:
 1. The adaptation time of the simulated inlets is longer than 100 years
 2. The inlets may never reach equilibrium due to missing or incorrect processes necessary for a stable equilibrium
- Significantly different results were obtained for different sediment transport capacity formula

Discussion and Conclusions

- Model computational times were reasonable
 - ▶ 100 years in about 7-10 days on a PC
- Model stability was very reasonable
- Cross-sectional areas were generally over-predicted
- Ebb and flood shoal morphologies and evolution were reasonable
- Comparison to the Escoffier curves were reasonable

Future Work

- Multiple grain sizes
 - ▶ Reduce channel erosion
 - ▶ Help reach dynamic equilibrium faster
- Dynamic roughness
 - ▶ Function of the bed gradation and bedforms
- Bank erosion feature
- Influence of jetties, asymmetric bays, and dredging
- Inlet infilling and closure?





Thank you

Questions?

